

## Chapter 5

### Discussion/Conclusions

#### Introduction

Cove Point fits Escoffier's (1954) model of a traveling foreland, but the geomorphic evidence indicates that the process of beach ridge progradation is episodic and variable rather than a constant, uniform process. Historic evidence and a comparative analysis of Flag Ponds indicates that phenomena such as erosion reversals and lateral spit growth are a common and integral part of foreland development. Prehistoric migration rates (e.g. 1.3 meters per year) differ from the migration rates derived through shoreline change analysis (e.g. 0.7 meters per year), indicating that sea level rise and especially bathymetry are playing an important role in recent advancement and morphology. The long-term morphology as indicated by vibracoring and radiocarbon dating reveals a linear pattern of migration which is driven primarily by a constant sediment supply regime and sea level rise. The beach ridges which make up the foreland are episodic in their constructional mechanism, although it is not clear what drives this cyclic behavior.

Cove Point, initial development.

The cusped foreland which is Cove Point probably did not exist more than 3000 years ago. Relative sea level was approximately 6 meters below its present level and rising at a rate of about 2 meters per 1000 years, (Peltier and Jiang, 1997). As the transgression continued, the coastline achieved its approximate present-day configuration. The littoral cells which Downs (1993) defined began to establish their patterns. Cove Point probably began as a small projection on the coast of Calvert County north of today's cape, possibly at a resistant ironstone headland such as Rocky Point (see Figure 51).

Rocky Point or a similar feature interrupted the longshore sediment transport regime, causing the development of a recurved spit and eventually a local reversal in the normal southerly drift as the cusped feature developed. Sediments accumulated in the form of spits which welded themselves onto the south shore of the cape to form beach ridges. At approximately 2500 years ago, Cove Point probably looked much like today's Flag Ponds. This process of accretion was rapid in the early stages of foreland development, with slower rates of erosion on the northern flank. This can be attributed to two factors. First, the young cape had not grown very far distally, and so was still in shallow water. Second, the rate of sea level rise was approximately 2 meters per 1000 years, much slower than present day rates (about 3.6 mm per year in the Chesapeake Bay due to a combination of glacial isostatic adjustment and eustatic sea level rise; Peltier and Jiang, 1997; Douglas, 1991).

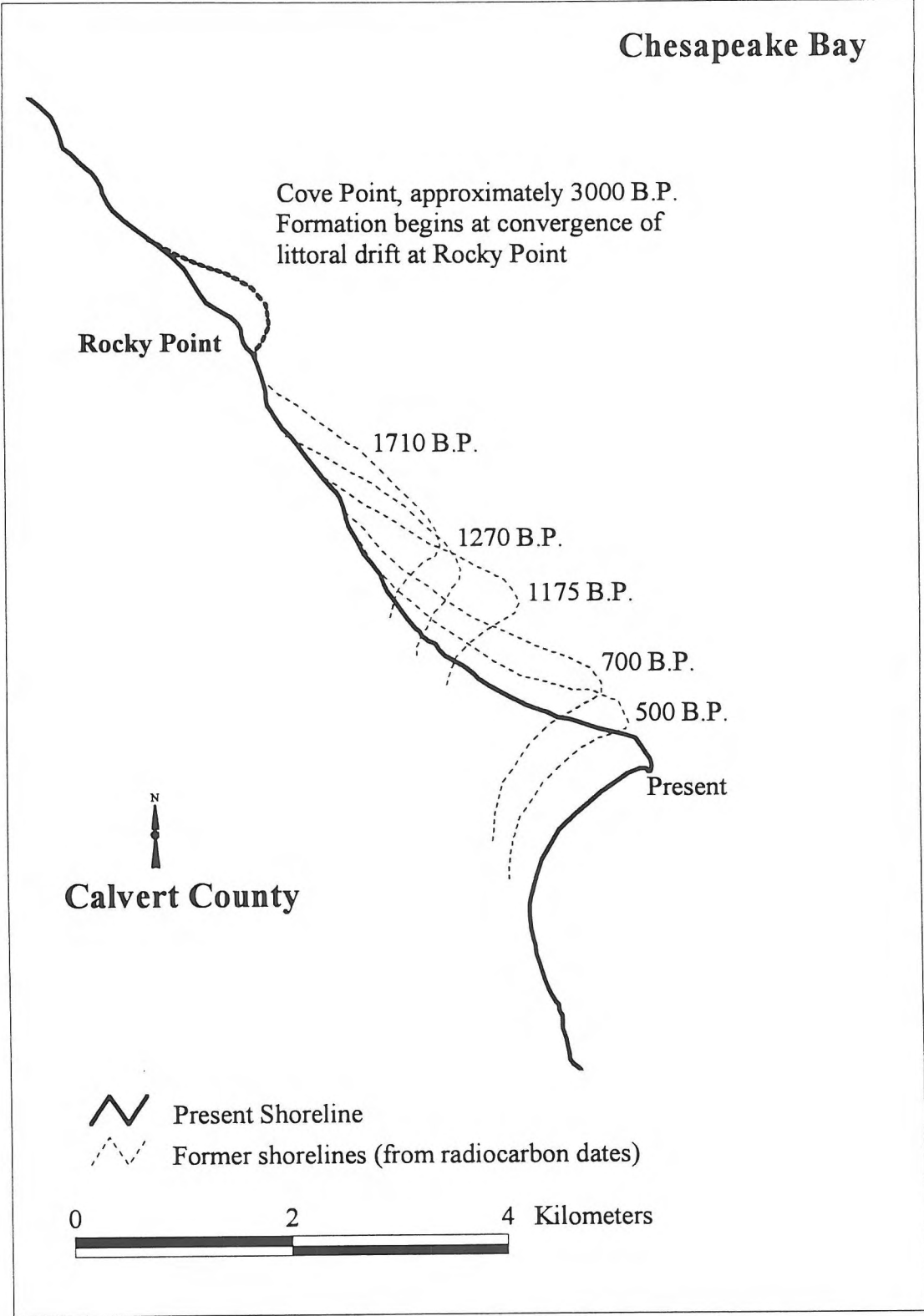


Figure 51. Cove Point initial inception and subsequent migration.

### Cove Point, prehistoric development

Cove Point continued its migration south as seen in Figure 51, with accretion taking place more rapidly on the southern flank than the erosion rate of the northern flank. This rate of migration over the last 1760 years has been 1.3 meters per year, a rate which is comparable to the present day migration of Flag Ponds (1.6 meters per year). Flag Ponds is currently developing in shallow water, much as the early Cove Point did, and an analysis of aerial photographs for the Flag Ponds region reveals the recent construction of cat-eye ponds. These ponds can also be seen in aerial photographs of Cove Point. They were probably formed when spits wrapped around the front of the foreland, cutting off areas of deeper water. These ponds quickly turned to freshwater with rainfall input, and a freshwater marsh began to grow. No saltwater peats were cored, indicating that the barrier strand which forms the Chesapeake Bay border of Cove Point remains fairly intact during the migration of the cape.

At approximately 700 years ago (according to radiocarbon dating), the growth of Cove Point reached a threshold. Spits no longer curled around the front end to cut off deeper water and create cat-eye ponds as at the present day Flag Ponds. This process formed ridges at a spacing of 30 meters at an average interval of 31 years. Instead, the beach ridges of Cove Point were constructed in the foreshore and welded directly onto the existing beachface. This process is more time consuming, encompassing a mean interval of 35 years, and results in a closer spacing of the beach ridges (25 meters). The long construction interval and

the apparent berm sequence found in the cores suggests that these ridges formed through a continual process of swash deposition as described by Tanner and Stapor (1972). This change in the constructional mechanism of the foreland is governed by the underlying Miocene surface of Cove Point which was previously the bottom of the Chesapeake Bay. This platform upon which the cape rests dips offshore gradually to the ancestral Susquehanna River channel. As the cape grew distally from the western shore of the Bay, it advanced into this deeper water. The Bay just south of the advancing flank was now at or below the depth of closure, limiting the movement of sediment for ridge construction. Also, the deeper this water becomes, the more sediment is required for each individual ridge. Hence there is an overall thickening from north to south in the wedge of littoral sediments as seen in the vibracores.

This increase in the amount of sediment required for each ridge is compounded by sea level rise. Sea level has risen approximately 2.0 meters during the last 1000 years (Peltier and Jiang, 1997), again increasing the water depth offshore of the advancing cape. This rising sea level produces the apparent tilting of the beach ridge plain (Figure 52). This tilting is not tectonic; it is due to the fact that the elevation of each ridge is controlled by the water level conditions present during the construction period. Thus, the older ridges are built at a lower sea level, while the higher ridges are built more recently. The oldest ridges have been completely drowned and covered by the vertically accreting marsh. This marsh, although not tidal, is controlled over the long-term by a rising water table

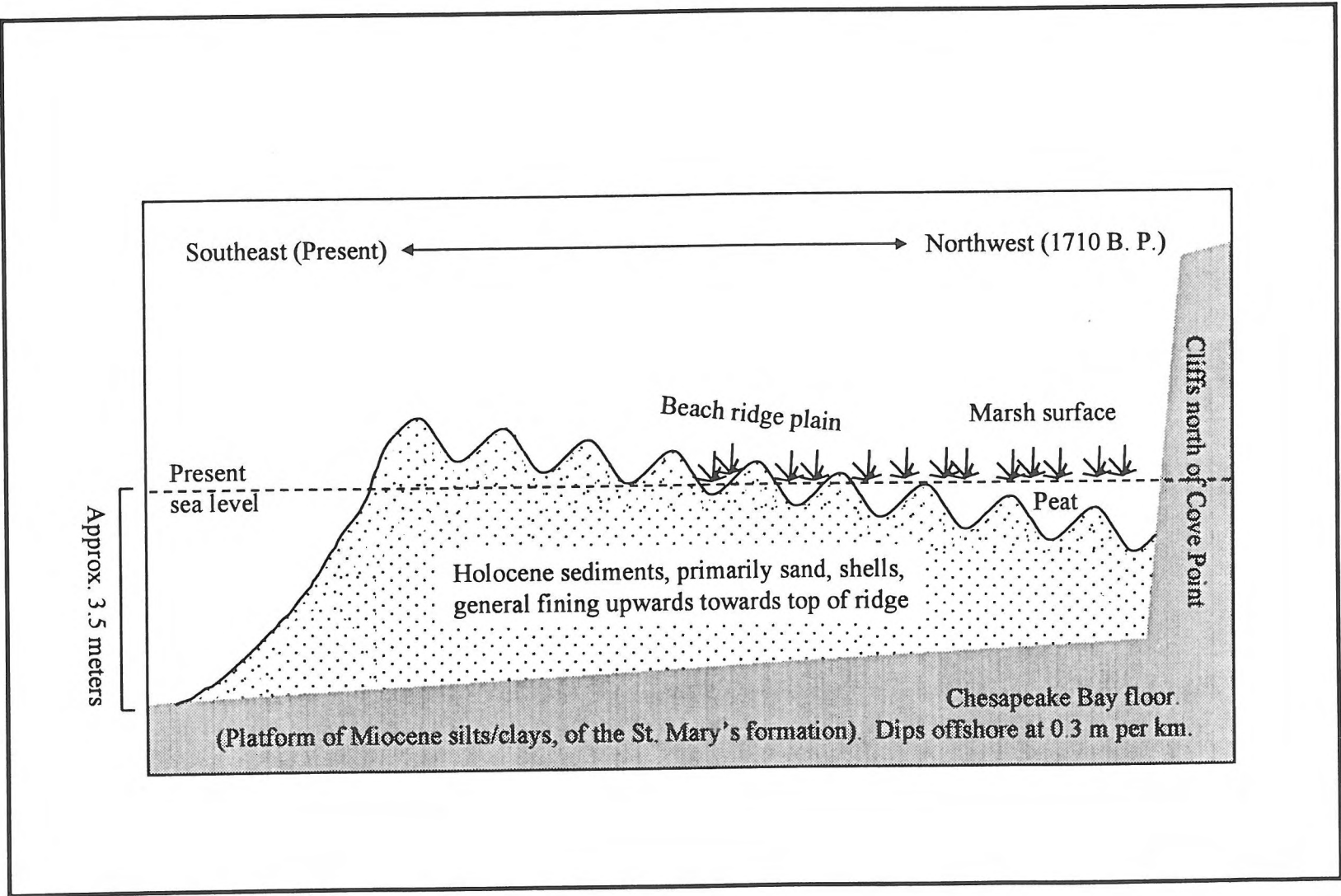


Figure 52. Diagram of vibrocore transect cross-section stratigraphy

which is controlled by sea level. Over the last three thousand years the marsh has been able to build on top of its own detritus, maintaining an elevation with respect to sea level. Interior pond expansion between 1938 and 1990 are indicators that the recent acceleration in relative sea level rise will interrupt this natural process.

#### Cove Point, historic development

Historic shoreline records indicate that Cove Point has been shrinking for at least the last 100 years. The northern flank is eroding at 1.0 meters per year while the southern flank accretes at 0.7 meters per year. This has led to a reduction in the subaerial size of the cape feature. This is partially controlled by the offshore bathymetry which is limiting the growth of the foreland as it advances into deeper water. It is also a product of the change in the rate of sea level rise over the last century. Sea level was rising at approximately 2.0 mm per year in the Chesapeake Bay prior to 1850 (Peltier and Jiang, 1997). Since that time, sea level has risen at a rate of 3.6 mm per year in the Bay (Douglas, 1991). This recent rise in sea level is the sum of the global eustatic rate (1.8 mm per year) and regional subsidence (Douglas, 1997; Peltier and Jiang, 1997). This rise in water level means a reduction in accretion rates and an increase in erosion rates for Cove Point as well as a continued deepening of the waters adjacent to the advancing flank of the foreland.

It appears Cove Point has reached and crossed a threshold between overall accretion and erosion. Previous Holocene migration occurred at a rate of 1.3

meters per year; it is now only 0.7 meters per year. Residents of Cove Point will see little accretion along the southern shore of Cove Point while the interior marsh continues to convert upland into freshwater marsh, encroaching on residential properties. It is probable that the shoreline engineering structures which protect the lighthouse are having some affect on this process, further aiding the drop in accretion as sediment is shunted offshore and lost to the sand-sharing system. There is an arc of erosion just south of the lighthouse, which represents the limit of influence of the seawall and groins. As such, their impact on the cape's morphology is not large. However, it is possible that this arc of influence will extend further south and west as foreland migration continues and sediment eroded from the northern flank continues to move offshore. It is also likely that there is lag time built into the system which is masking the effect of these structures. A beach ridge presently requires 35 years to build; the sediment which is accreting onto the southern flank may have been stored in the shoal before the groins were built. As such, the effect of these groins may not be noticeable for years to come.

The beach ridges which make up Cove Point are presently being constructed at a rate of 35 years per ridge with an average spacing of 25 meters. Ridges in the older part of the foreland probably emerged as bars, cutting off deeper water (a process which can be inferred through an analysis of core #7). This process is similar to the mechanism described by Curray et al. (1967). As the foreland advanced into deeper water, ridges welded directly onto the beachface rather than emerging as offshore bars. These ridges are driven by a cyclical



process; a certain threshold within the geomorphic system must be reached before a ridge is constructed. It is likely that this threshold is reached by the accumulation of sediment in the shoal at the tip of Cove Point. Once enough sediment is present, a high energy event such as a northeaster moves this sediment, distributing it alongshore so that it can be reworked as nearshore bars and eventually beach ridges (Figure 53).

It was not possible to draw any conclusions about the topography of the underlying Miocene surface. This surface may have been dissected by riverine channels which ran across the Susquehanna floodplain prior to the most recent transgression (Kerhin, 1997, pers. comm.). This is supported by the unusually thick littoral sequence found in core #5; however, it is not possible to map any undulations of this surface due to the poor sampling resolution. Errors in vibracore compaction/correction and elevational positioning also makes it difficult to draw solid conclusions about the dip of the Miocene strata. However, when the Miocene contact is plotted against depth, there does appear to be a slight dip to the south (approximately 0.3 meters per kilometer along the vibracore transect). A dip of 1.4 to 2.0 meters per kilometer is indicated by Kidwell (1997); this discrepancy can again be explained by the coarse sampling resolution and possible errors when accounting for vibracore compaction.

Cove Point is a unique feature which has developed under a set of conditions which are particular to that stretch of the Calvert County coastline. Dominant wave attack from the north drives erosion; frequent wave action from

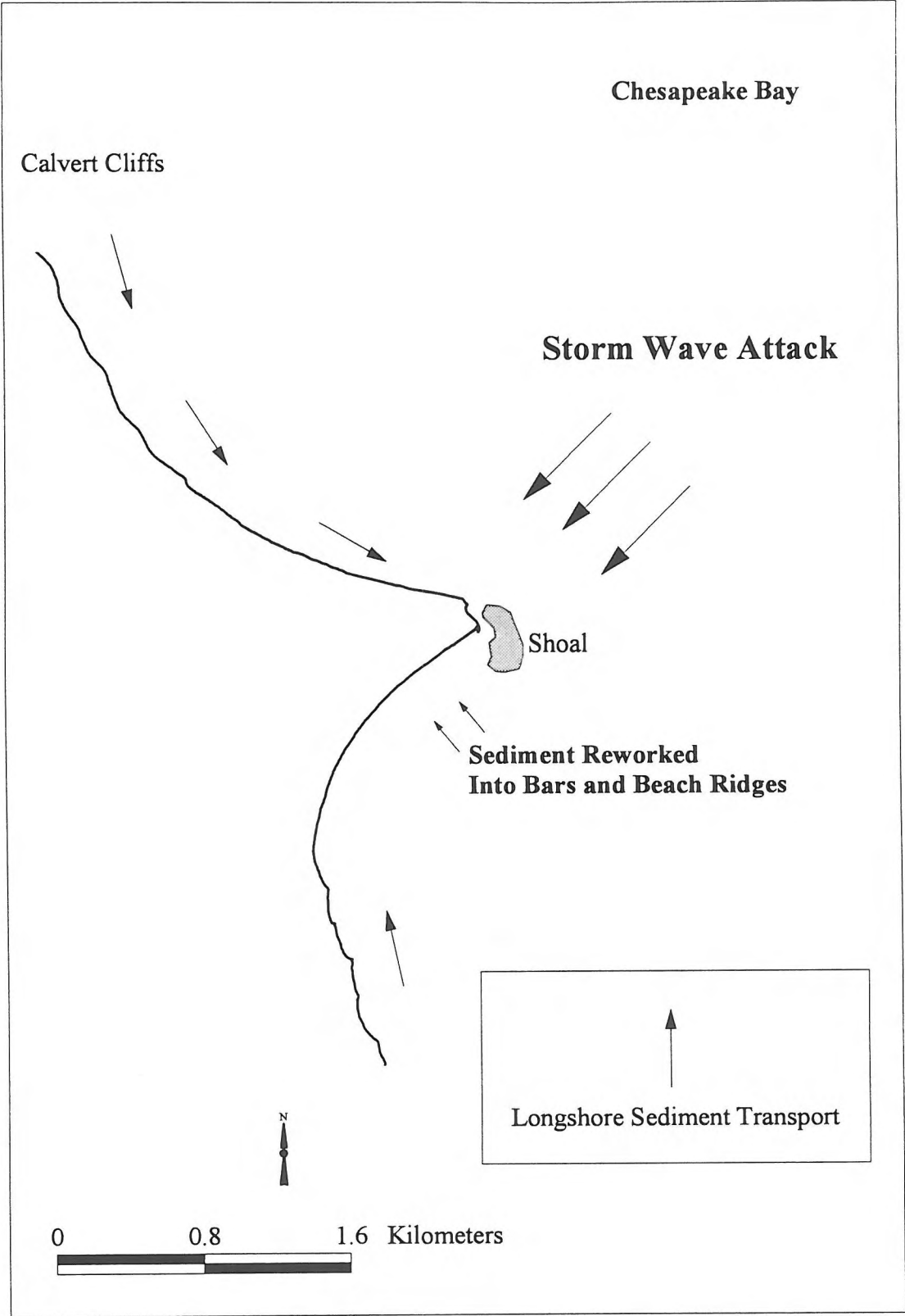


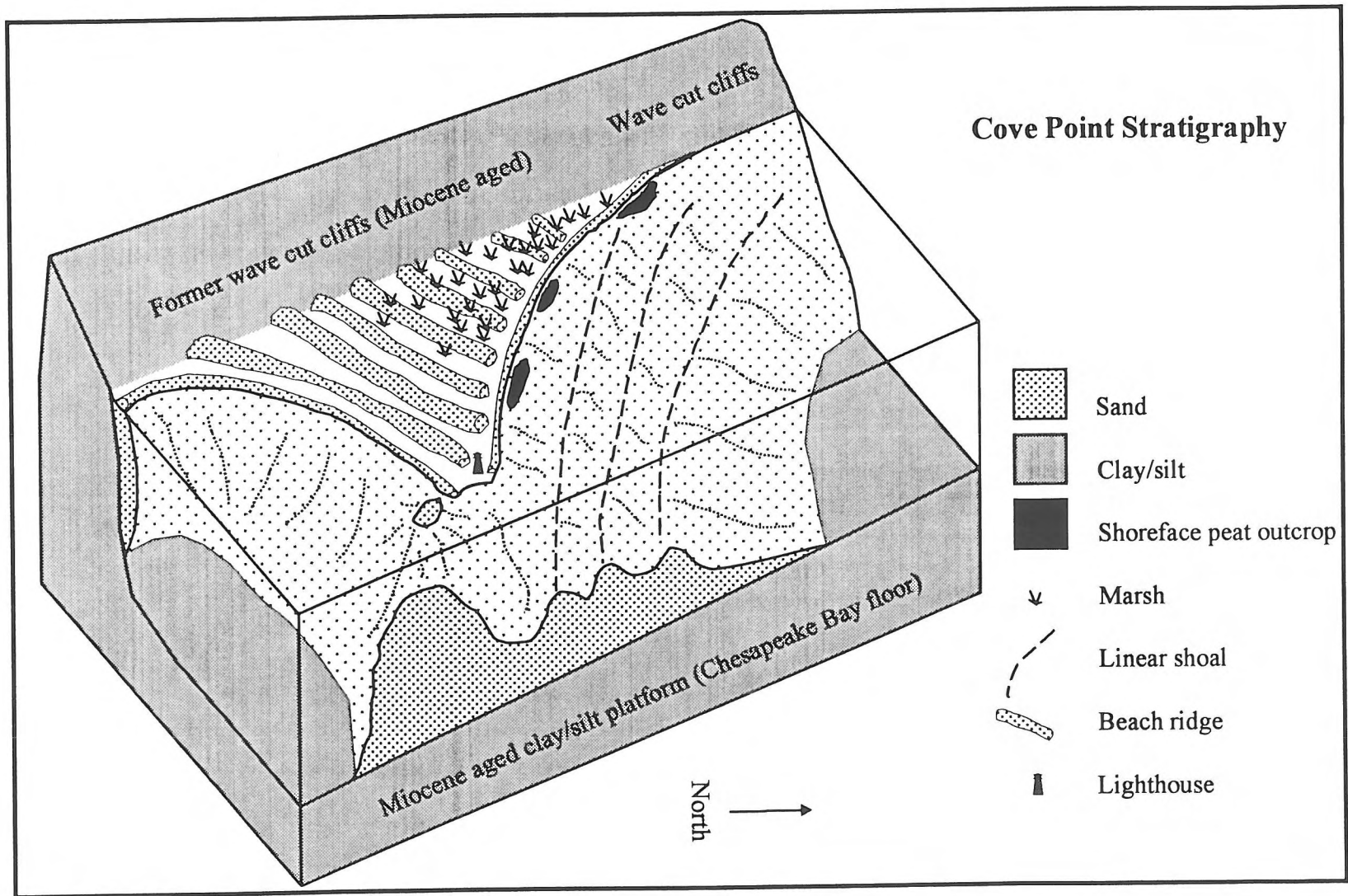
Figure 53. Map of Cove Point showing drift and ridge construction.

the south causes accretion (Escoffier, 1954). The restricting eastern shore of the Chesapeake Bay limits fetch and hence wave energy from the west, allowing distal extension of the cape feature into deeper water (Figure 54). It is interesting to compare Cove Point with other traveling forelands. Cape Hatteras and Flag Ponds have similar construction mechanisms (e.g. one flank progrades through the growth of beach ridges, while another flank is eroded), although in each case they are operating at different temporal and spatial scales. Cape Hatteras is several orders of magnitude larger than Cove Point, while Flag Ponds is an order of magnitude smaller (Figure 55).

#### The future of Cove Point

The present day morphology of Cove Point is governed primarily by the bathymetry off the southern flank of the cape which is limiting further growth. The shore engineering structures at the tip of the cape are also having an effect on morphology, possibly shunting sediment offshore before it can accrete onto the southern flank. Sea level rise is the least of these three factors; it has controlled the long-term elevational construction of the beach ridge plain and will continue to do so in the future; at present, it is causing increased rates of erosion on the northern flank and slower rates of accretion on the southern flank. In 50 years, Cove Point will look very similar as the same general trends of erosion and accretion continue; there will be an overall reduction in size of the subaerial cape. The tip of the spit may also thin down if the minor erosion noted on the southern

Figure 54. Block diagram of Cove Point stratigraphy.



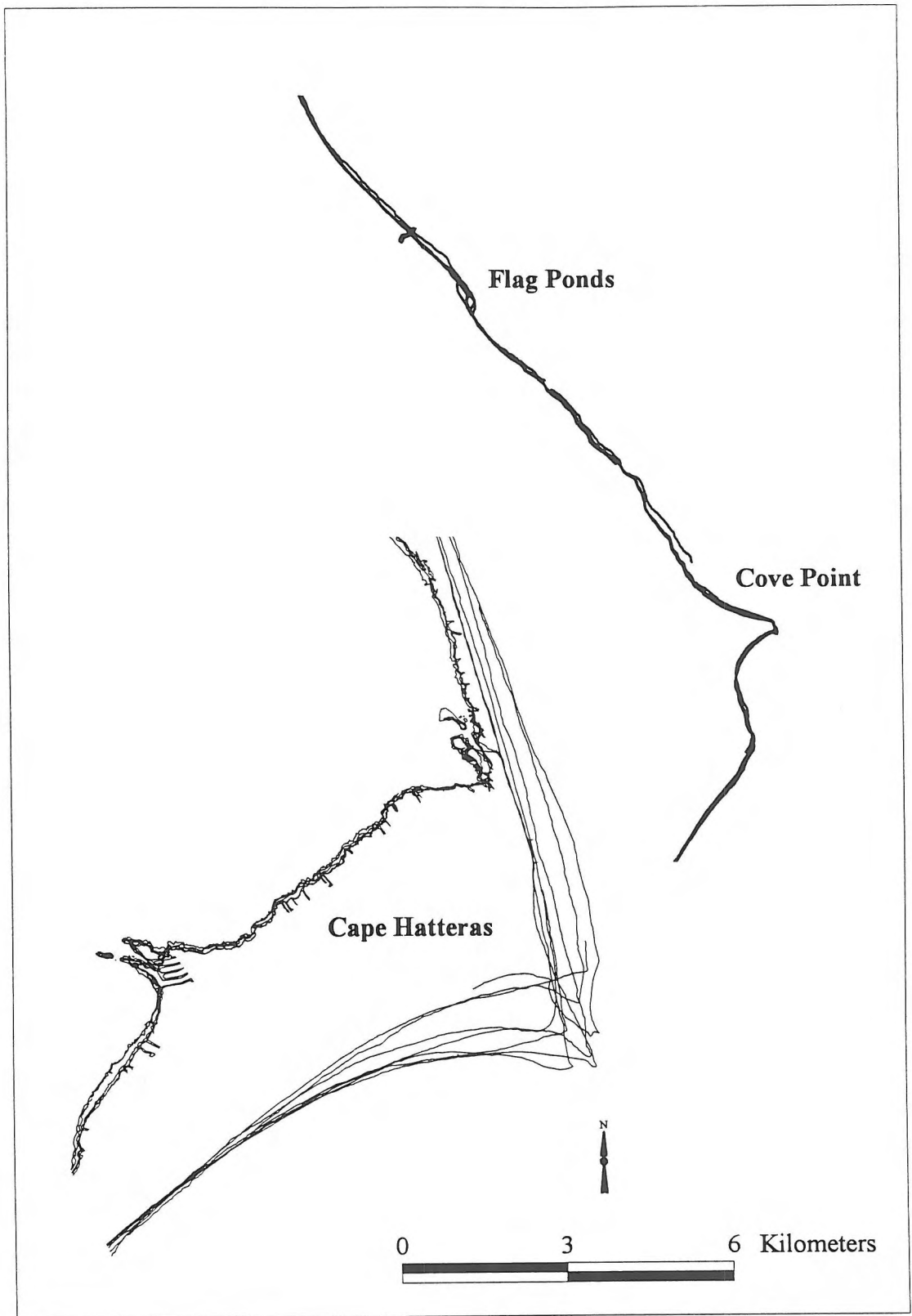


Figure 55. Comparison of cusped forelands.

flank (just southwest of the tip) continues. Sea level rise will continue to cause the interior marsh to encroach on the residential community located on the youngest part of the cape feature.

#### Suggestions for further research

There is a great deal which remains to be discovered about the geomorphic development of Cove Point. Although beach ridge stratigraphy is investigated in this study, it was not possible to fully reconstruct the advance of these ridges for the entire foreland. This was particularly true for the oldest parts of the beach ridge plain which are covered by marsh. A comprehensive study of beach ridge topography would be very useful. Such surveys could be accomplished through the use of ground-penetrating radar. This area is ideal for such a project; this radar would easily penetrate the marsh and reveal the detailed topography of the beach ridge plain.

Other potential work includes more radiocarbon dates of vibracore organic materials, a more intensive coring program (possibly with other methods such as a Dutch auger), and creating a more temporally dense shoreline record database by incorporating other vertical aerial photographs. Other useful work would be a comparative study of Flag Ponds using vibracoring and radiocarbon dating. Flag Ponds is an incipient cusped foreland probably similar to what Cove Point must have been a few thousand years ago. The conditions which shaped Cove Point are

in an earlier stage of operation at Flag Ponds; a comparative study here would lend insight to the early evolution of Cove Point.

## Appendix A

### Vibracore Logs

These logs are records of the vibracores taken at Cove Point in the summer of 1996. They have been corrected for compaction; as such, depths do not match the depths observed in the core photographs (Appendix B). Logs are arranged with the topmost part of the core starting in the top left corner of the page. At the end of two meters, the core log starts again on the right side of the page. Because all the cores were uniform after 4 meters (they all sampled the St. Mary's formation if they reached that far) it was decided that only the first 4 meters would be shown. Each unit is described; carbon content test results are noted in the middle of the graphic core display. The legend for the cores requires a few notes of explanation: The division between fine and coarse grained sands is 1.0 phi ( $\emptyset$ ). Those sediments which were classified as sand with shells were typically coarse grained; thus it was not necessary to have a fine grained sand with shells category. The clay and silt category was used almost exclusively for the Miocene deposits which appeared in many of the cores. This was the only type of clay/silt observed in the stratigraphic logs. Variable is classified only where the mixing of layers is too dense to permit detailed explanations. Grain sizes are given in phi ( $\emptyset$ ) and millimeters as follows: 2.0 = 0.25 mm; 1.5 = 0.357 mm; 1.0 = 0.500 mm; 0.5 = 0.707 mm; 0.0 = 1.000 mm; -0.5 = 1.414 mm; -1.0 = 2.000 mm.



